

High resolution atmospheric modelling on the Tibetan Plateau : How to understand a system with sparse observations?

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The Tibetan Plateau is a key environment for global change

- The Tibetan Plateau (Fig. 1) is located north of the Himalaya, it is the world's largest mountain highland and has a mean elevation of more than 4500 m with large parts covered by *Kobresia* pastures.
- Important for regional climate as largest glaciated area outside the poles (3rd pole) (i.e. Qui, 2008), source region of East-Asian river systems with > 1.5 billion people living in their catchments (water towers for humanity) (i.e. Immerzeel, 2010) and modifier of monsoon circulation (Boos & Kuang, 2010).
- The surface energy balance (Fig. 2) is very variable and depends on

local influences of topography, cloud cover and soil moisture.
•The surface energy balance is important for weather and climate.

→ Crucial to get it right

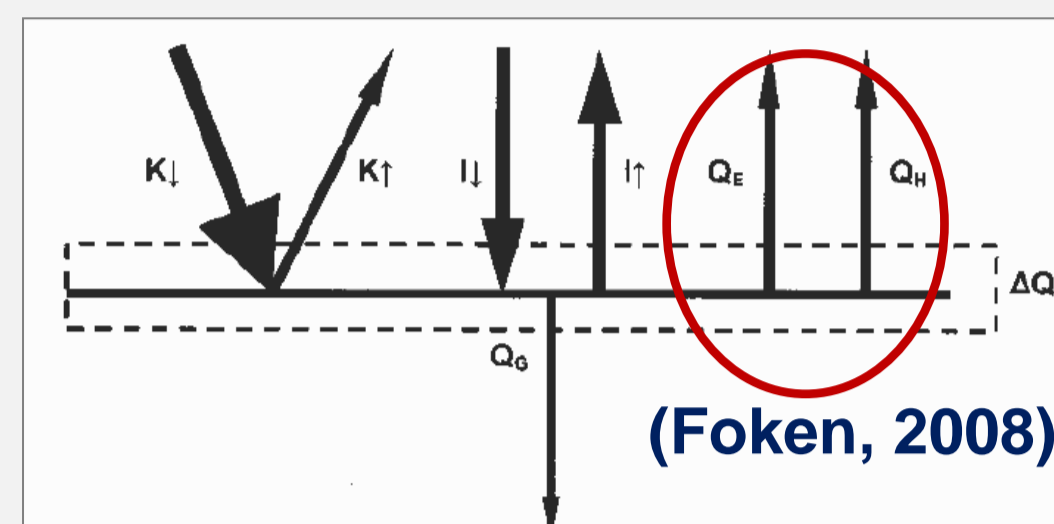
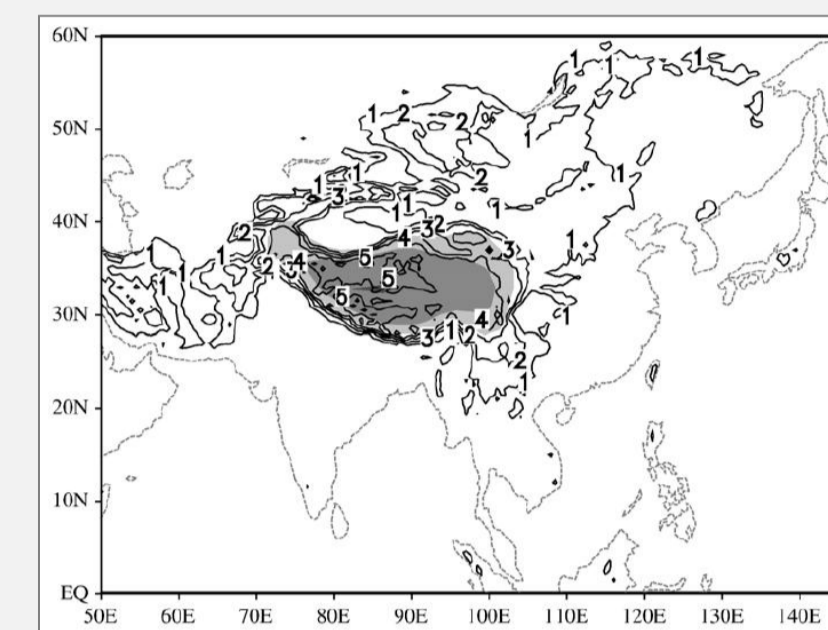


Fig. 2: Concept of surface energy balance with shortwave (K), long-wave (l) radiation and its distribution into turbulent fluxes of heat (Q_h) and water vapor (Q_E) being dependent on surface conditions (Foken, 2008)

Fig. 1: The Tibetan Plateau and its elevation (shaded - Cui et al., 2006; Glob. Plan. Change).



Sparse Observations = Uncertain Models

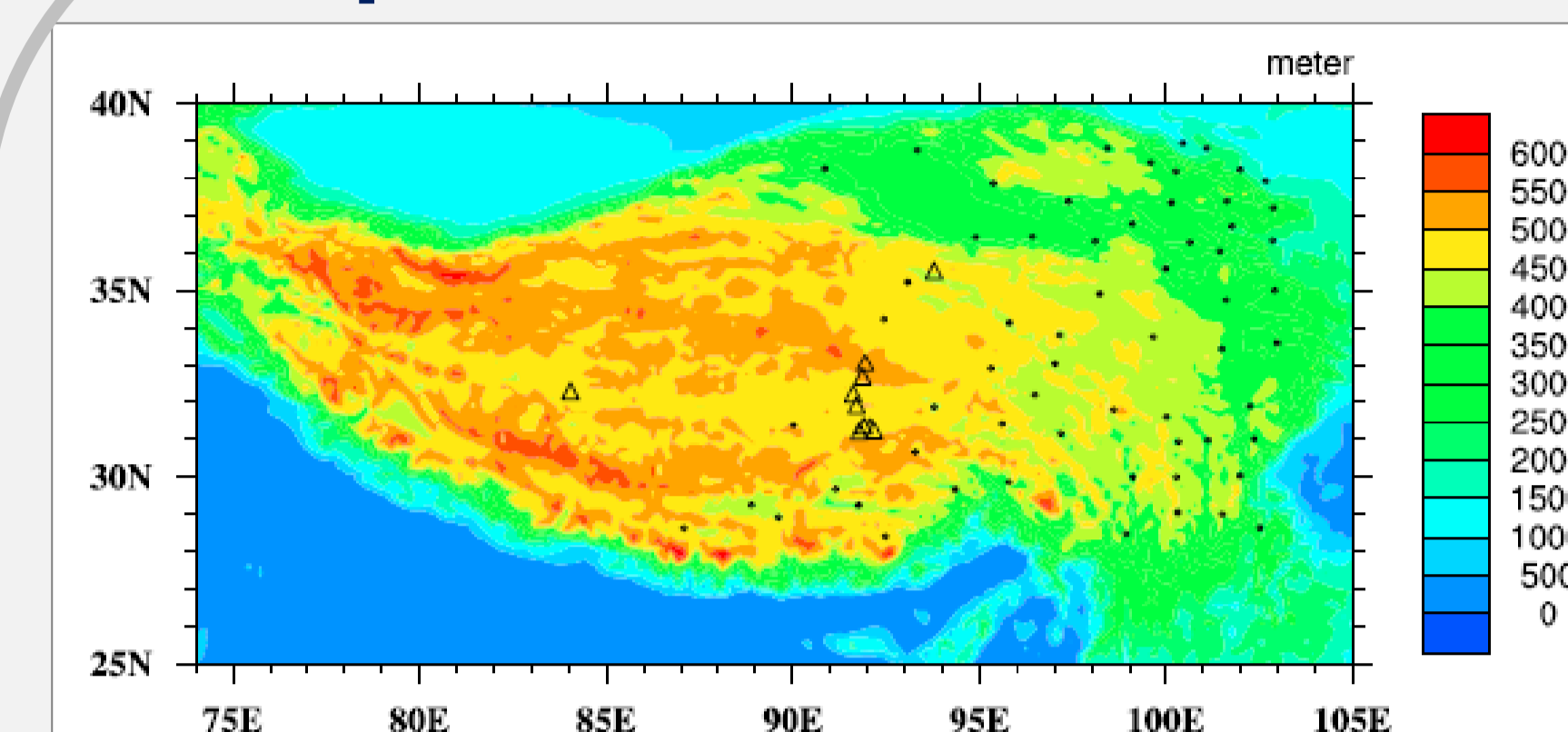


Fig. 3: Topography and locations of weather stations (CMA: dots, CAMP: triangles) on Tibetan Plateau (Wang & Zhang, 2012; JGR)

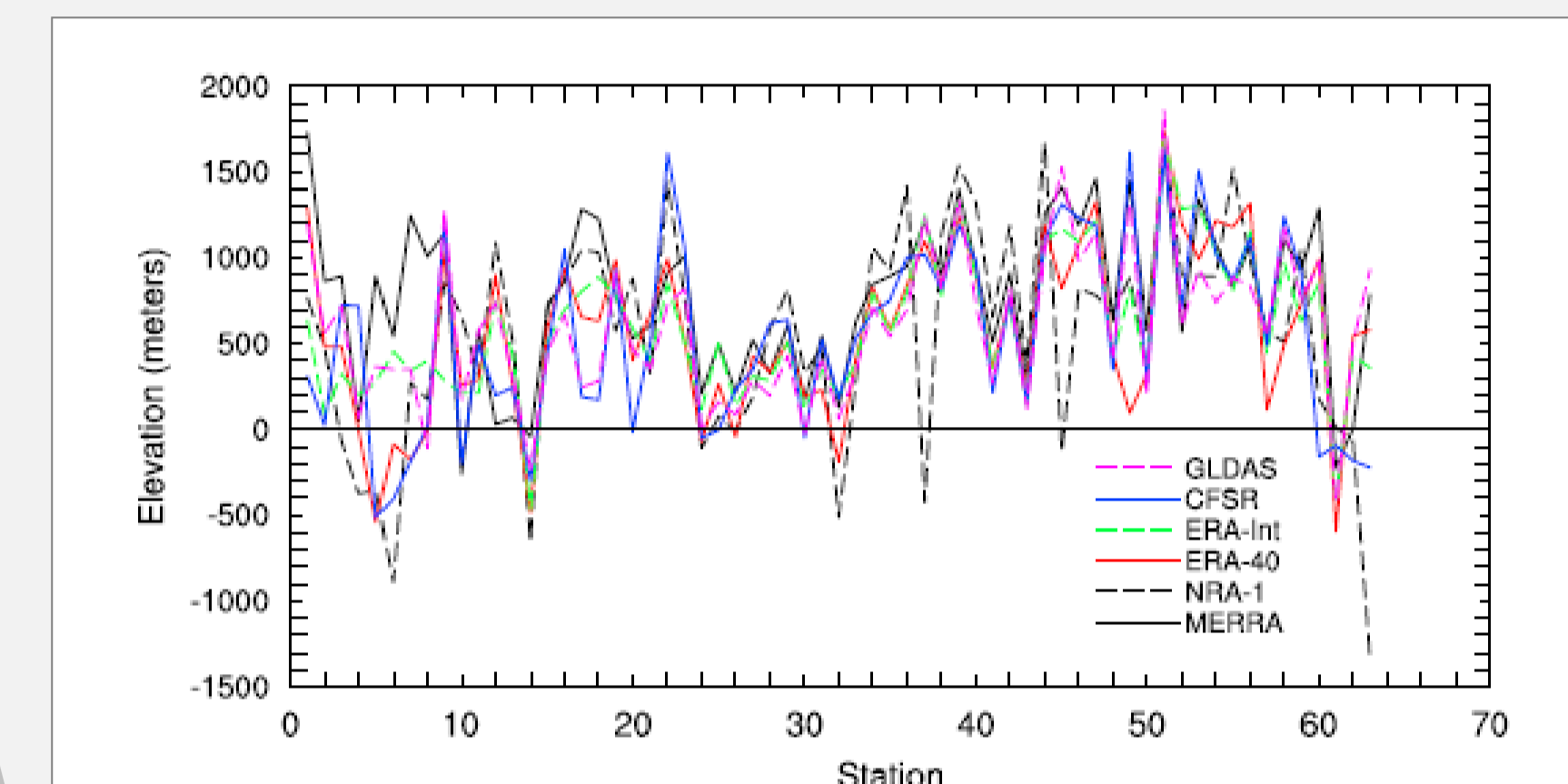


Fig. 4: Surface elevation differences between reanalysis grid cells and station locations. (Wang & Zhang, 2012; JGR)

- Tibetan Plateau is a remote area with very complex terrain and variable surface conditions.
- This means that access and running high quality, continuous measurements is difficult and station locations are often not chosen for representativeness of an area, but facility of access.
- Weather stations are concentrated in the "populated" and lower eastern TP (Fig. 3).
- Stations are usually at lower elevation than corresponding reanalysis grid cell (Fig. 4) and no permanent stations above 4800 m (Maussion, 2011)
- Reanalysis and satellite data not reliable and often with coarse resolution like degree scale ($1^\circ \approx 100\text{km}$)

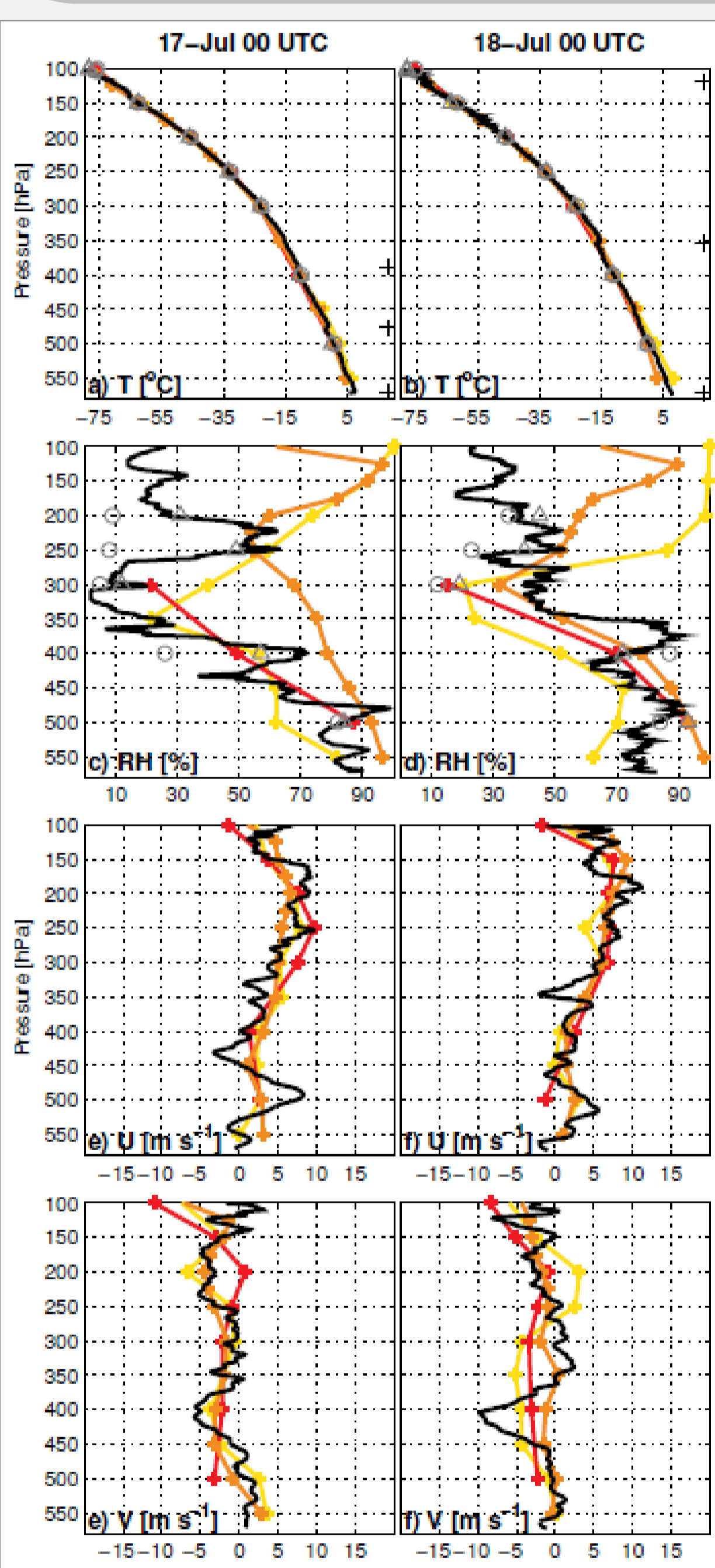


Fig. 7: Atmospheric profiles for Nam Co station on 17+18.7.2012, from radiosondes, NCEP-I, ERA-Int, GFS-FNL (Gerken et al., 2013b, subm. JGR)

- There is a large uncertainty for model boundary conditions: Both for the atmosphere (Fig. 7) and for the surface.
- We have shown that our modelling approach is capable of reproducing the most important interactions between the surface and the atmosphere and atmospheric dynamics (Fig. 8).
- Using different atmospheric conditions: Radiosondes and gridded data products as commonly used for climate research, the model produces very different convection developments (e.g. thunderstorms) and amounts of rain (Fig. 9)
- Thus, through cloud surface interactions, the surface energy balance is quite different (Fig. 10) → Large uncertainty
- Typically not represented in coarser resolution models as clouds are not resolved and profiles are taken from stock

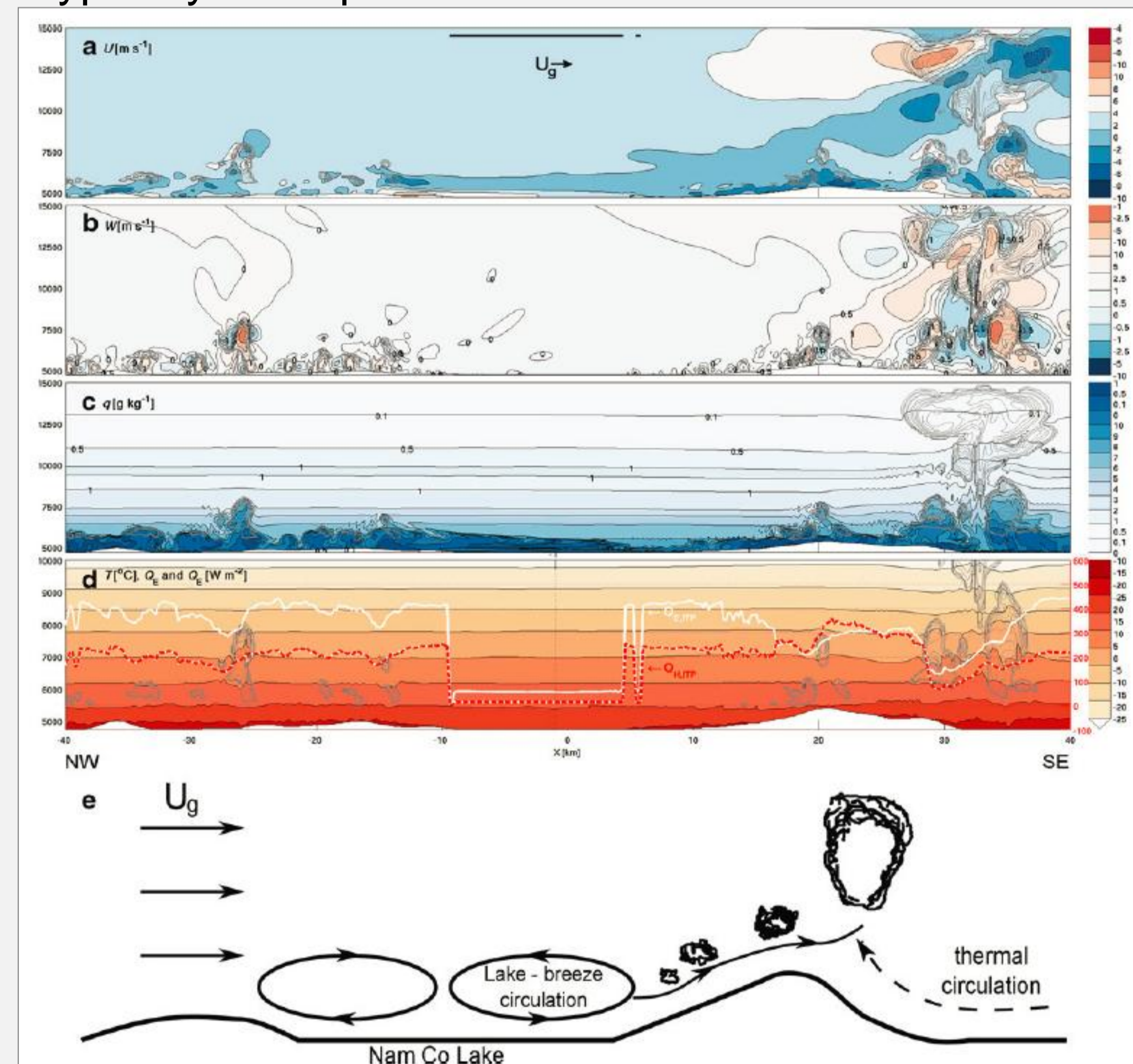


Fig. 8: Illustration of convection development at Nam Co Lake for 6.8.2009 (Gerken et al., 2013a, Theor. Appl. Clim.)

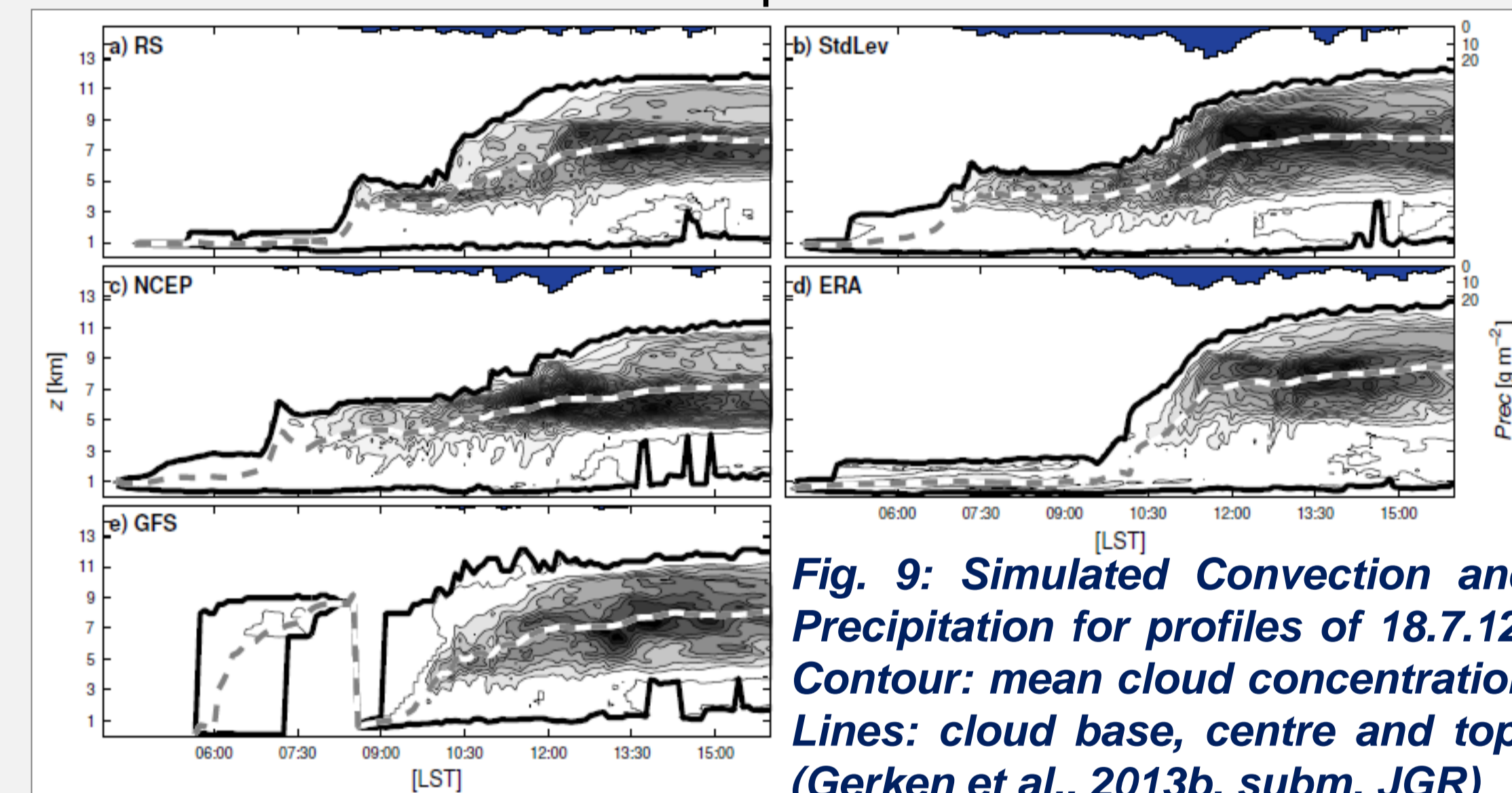


Fig. 9: Simulated Convection and Precipitation for profiles of 18.7.12. Contour: mean cloud concentration Lines: cloud base, centre and top: (Gerken et al., 2013b, subm. JGR)

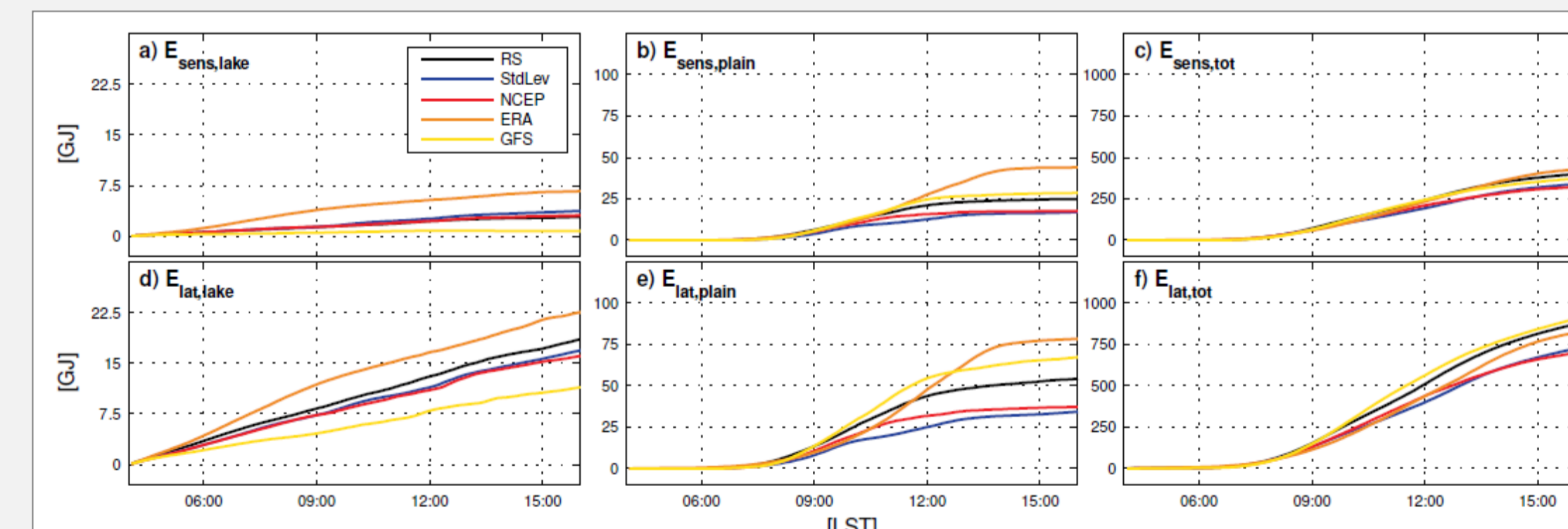


Fig. 10: Simulated surface energy transfer for lake, plain area and total domain for profiles of 18.7.12 (Gerken et al., 2013b, subm. JGR)

High Resolution Modelling at Nam Co Lake

- Mesoscale models typically operate on ~10 km resolution, which is too coarse to adequately represent topography. Input information is even coarser.
- High resolution studies in order to estimate behavior and sensitivities of system.
- Our area of interest is the Nam Co Lake basin (Fig. 5) 4730m above sea level and surrounded by mountain chains up to 7000 m high.
- We use the Active Tracer High-resolution Atmospheric Model – ATHAM (Herzog et al., 1998 ; Oberhuber et al., 2008) for 2D simulations with 200 m horizontal resolution .
- Simulation of daily cycles: Surface fluxes driven by solar radiation lead to the evolution of fair weather clouds to convection (tall clouds with precipitation) (Fig. 6).
- Investigation of the influence of atmospheric and surface conditions.

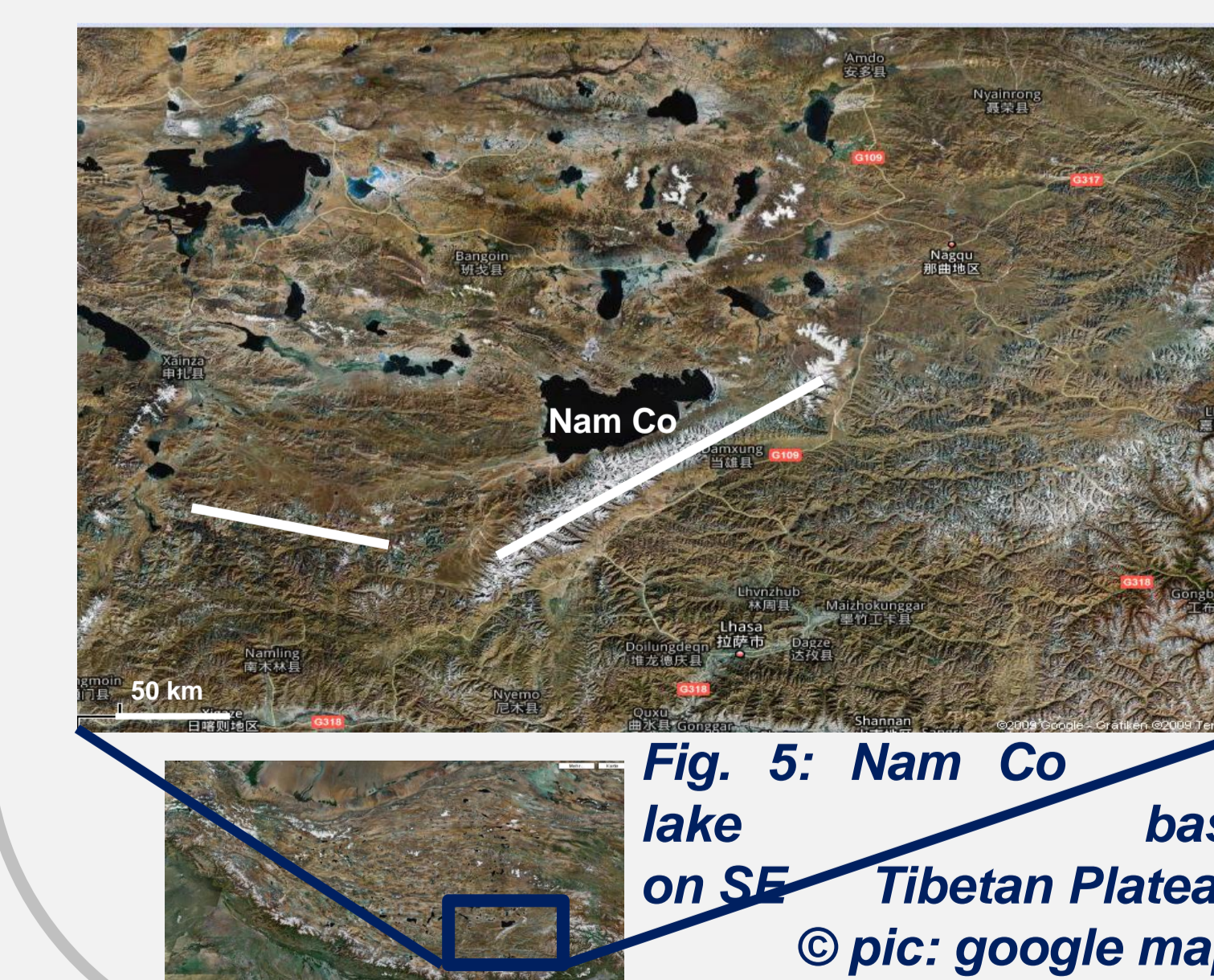


Fig. 5: Nam Co lake basin on SE Tibetan Plateau © pic: google maps

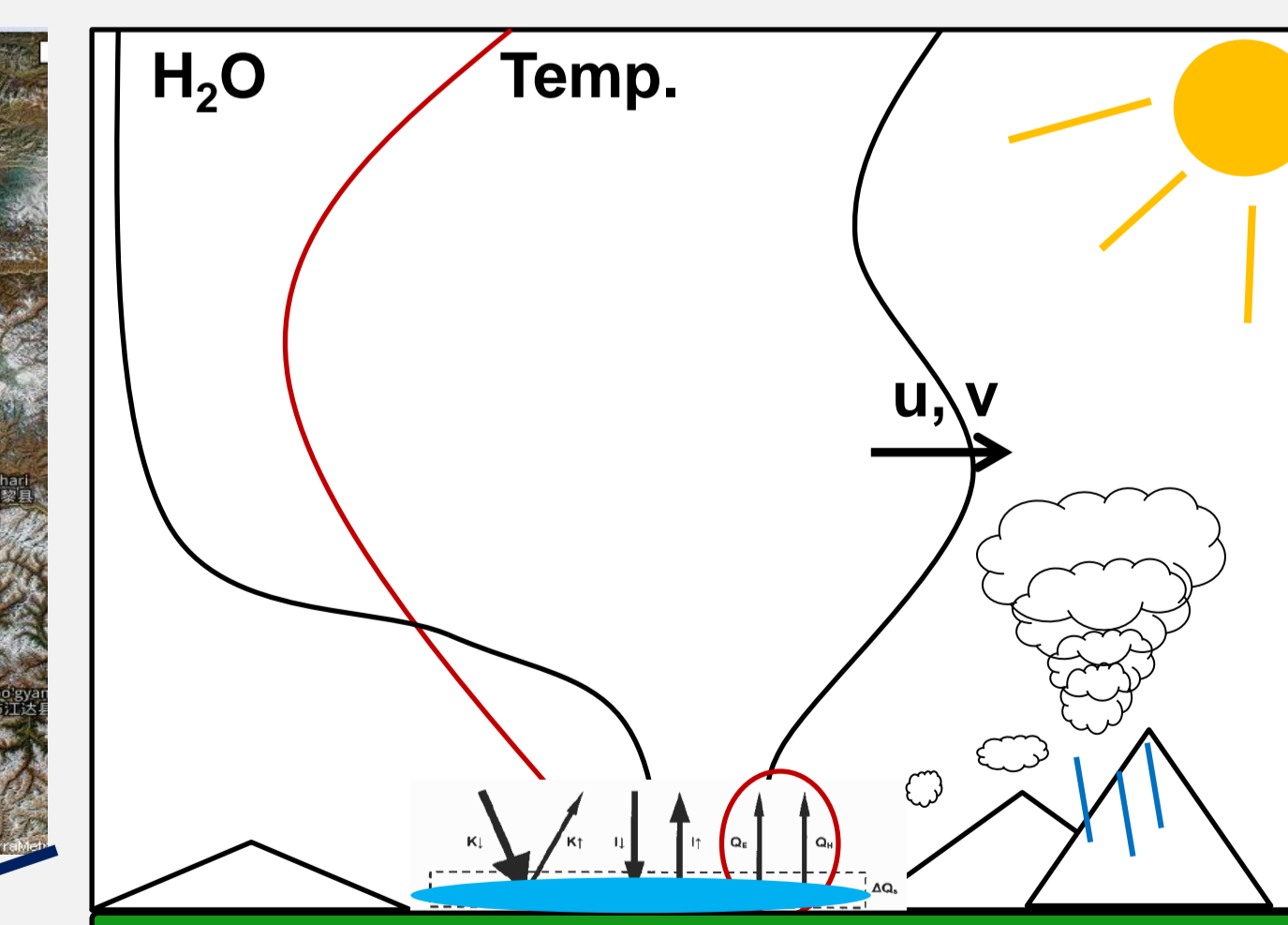


Fig. 6: Schematic of lake, pasture, mountain system: Nam Co Lake and evolution from small clouds to convection.

- Large uncertainty of measurements due to remote location and complex terrain.
- "Normal" models are too coarse to resolve topography, clouds and use very uncertain input data.
- This leads to systematic errors in simulated surface energy balance and weather.
- Thus, novel approaches are needed to investigate uncertainties for climate.

References

- Qui, J.: China: The third pole, *Nature*, 454, 393–396, 2008.
- Immerzeel, W. W., van Beek, L. P. H. and Bierkens, M. F. P.: Climate Change Will Affect the Asian Water Towers, *Science*, 328(5984), 1382–1385, doi:10.1126/science.1183188, 2010.
- Boos, W. R. and Kuang, Z.: Dominant control of the South Asian monsoon by orographic insulation versus plateau heating, *Nature*, 463(7278), 218–222, doi:10.1038/nature08707, 2010.
- Wang, A. and Zeng, X.: Evaluation of multireanalysis products with in situ observations over the Tibetan Plateau, *J. Geophys. Res.*, 117(D5), D05102, doi:10.1029/2011JD016553, 2012.
- Maussion, F., Scherer, D., Finkelnburg, R., Richters, J., Yang, W. and Yao, T.: WRF simulation of a precipitation event over the Tibetan Plateau, China – an assessment using remote sensing and ground observations, *Hydrol. Earth Syst. Sci.*, 15(6), 1795–1817, doi:10.5194/hess-15-1795-2011, 2011.
- Herzog, M., Graf, H.-F., Textor, C. and Oberhuber, J. M.: The effect of phase changes of water on the development of volcanic plumes, *J. Volc. Geotherm. Res.*, 87(1-4), 55–74, doi:10.1016/S0377-0273(98)00100-0, 1998.
- Oberhuber, J. M., Herzog, M., Graf, H.-F. and Schwanke, K.: Volcanic plume simulation on large scales, *J. Volc. Geotherm. Res.*, 87(1-4), 29–53, doi:10.1016/S0377-0273(98)00099-7, 1998.
- Gerken, T., Biermann, T., Babel, W., Herzog, M., Ma, Y., Foken, T. and Graf, H.-F.: A modelling investigation into lake-breeze development and convection triggering in the Nam Co Lake basin, Tibetan Plateau, *Theor. Appl. Climatol.*, online first, doi:10.1007/s00704-013-0987-9, 2013.